



Cite this article: Proshutinsky A, Dukhovskoy D, Timmermans M-L, Krishfield R, Bamber JL. 2015 Arctic circulation regimes. *Phil. Trans. R. Soc. A* **373**: 20140160.
<http://dx.doi.org/10.1098/rsta.2014.0160>

Accepted: 13 July 2015

One contribution of 6 to a discussion meeting issue 'Arctic sea ice reduction: the evidence, models and impacts (Part 2)'.

Subject Areas:

oceanography, climatology, meteorology, atmospheric science

Keywords:

arctic climate variability, circulation regimes, freshwater and heat content

Author for correspondence:

Andrey Proshutinsky
e-mail: aproshutinsky@whoi.edu

Arctic circulation regimes

Andrey Proshutinsky¹, Dmitry Dukhovskoy²,
Mary-Louise Timmermans³, Richard Krishfield¹
and Jonathan L. Bamber⁴

¹Physical Oceanography Department, Woods Hole Oceanographic Institution, Woods Hole, MA, USA

²Center for Ocean-Atmospheric Prediction Studies, Florida State University, Tallahassee, FL, USA

³Department of Geology and Geophysics, Yale University, New Haven, CT, USA

⁴School of Geographical Sciences, University of Bristol, Bristol, UK

Between 1948 and 1996, mean annual environmental parameters in the Arctic experienced a well-pronounced decadal variability with two basic circulation patterns: cyclonic and anticyclonic alternating at 5 to 7 year intervals. During cyclonic regimes, low sea-level atmospheric pressure (SLP) dominated over the Arctic Ocean driving sea ice and the upper ocean counterclockwise; the Arctic atmosphere was relatively warm and humid, and freshwater flux from the Arctic Ocean towards the subarctic seas was intensified. By contrast, during anticyclonic circulation regimes, high SLP dominated driving sea ice and the upper ocean clockwise. Meanwhile, the atmosphere was cold and dry and the freshwater flux from the Arctic to the subarctic seas was reduced. Since 1997, however, the Arctic system has been under the influence of an anticyclonic circulation regime (17 years) with a set of environmental parameters that are atypical for this regime. We discuss a hypothesis explaining the causes and mechanisms regulating the intensity and duration of Arctic circulation regimes, and speculate how changes in freshwater fluxes from the Arctic Ocean and Greenland impact environmental conditions and interrupt their decadal variability.

1. Introduction

In this paper, we build on previous investigations of Arctic Ocean wind-driven circulation regimes [1–6] and

© 2015 The Authors. Published by the Royal Society under the terms of the Creative Commons Attribution License <http://creativecommons.org/licenses/by/4.0/>, which permits unrestricted use, provided the original author and source are credited.

their transformations under rapidly changing Arctic conditions, and extend them by focusing on identifying the causes and mechanisms responsible for decadal variability—in particular, on the deviation from climatological decadal variability since 1997. Atmospheric circulation, ice drift and wind-driven ocean motions are dynamically connected and responsible for significant changes in the state and variability of Arctic environmental parameters. The presence of relatively mobile sea ice, driven by winds and ocean currents, and regulating atmosphere/ocean interactions is a key factor influencing Arctic climate and changes on seasonal, interannual and decadal time scales. This has been recognized since Nansen's era [7] of Arctic exploration, and studied and discussed in numerous publications (e.g. [8–11]). In the following, we review past studies of Arctic decadal variability and circulation-regime definitions to introduce our hypothesis which may explain the breakdown of previously observed regular decadal variability.

(a) Arctic climate changes: 1948–2013

A representative suite of various Arctic climate data are available from approximately 1948; these observations have been summarized recently (e.g. [12,13]) to show that during the last two to three decades Arctic atmosphere and ocean temperatures have increased, sea ice volume and extent have decreased, permafrost has thawed, storminess has increased, sea level has risen and biological processes have become more complex and diverse [13].

We begin with an assessment of the 1948–2013 time series of Arctic environmental parameters, including sea ice extent, air temperature, sea-level atmospheric pressure (SLP), sea level, geostrophic wind, river discharge, Greenland freshwater flux and areal ice flux via Fram Strait (figure 1; for data sources, see table 1). We apply a 30-year running mean to all of the annual data (black lines in figure 1) to examine general trends in these parameters with the understanding that 'climate' is defined as weather averaged over a 30-year period (thick black lines in figure 1) following World Meteorological Organization and IPCC [17] recommendations. This averaging provides information on climate changes from approximately 1963 to 1998, and demonstrates that the Arctic climatological annual values of: (a) sea ice extent reduced by about 14%, (b) air temperature increased by 10% (from -12.55°C in 1963 to -11.30°C in 1998), (c) SLP reduced by approximately 0.7 hPa; (d) sea level rose by approximately 0.06 m; (e) geostrophic wind speed over the Arctic Ocean increased by about 15%; (f) Eurasian river runoff increased by 14%; (g) freshwater flux from Greenland has increased rapidly since the 1980s with acceleration after 1997 (present values are around 60% of the magnitude of Eurasian river discharge); and (h) areal ice flux via Fram Strait increased by 15% which must have had a significant influence on the total ice volume in the Arctic Ocean and specifically on the minimum of ice extent in September [18].

Changes in the atmospheric circulation between the climate state of 1948–1977 and that of 1984–2013 (centred, with 30-year mean averaging, on the years 1963 and 1998, respectively) are not visually dramatic (figure 2, upper panels) although there is a tendency to intensification of cyclonic atmospheric motion due to a reduction of SLP over the Arctic Ocean, and a reduction in area of the Arctic High centred over the Beaufort Gyre region (figure 2, left bottom panel). Mean geostrophic wind is intensified along the Norwegian coast (figure 2, right bottom panel) apparently increasing the Atlantic water inflow (and heat transport) to the Barents Sea [19]. At the same time, winds are stronger in the latter period in the Fram Strait region, increasing the sea-ice flux from the Arctic Ocean towards the Greenland Sea (e.g. [20]) and in the northern Canadian Archipelago straits, potentially increasing the ice and water flux from the Arctic Ocean towards Baffin Bay and the Labrador Sea [21].

The preceding analysis corroborates the general agreement among scientists that the Arctic has been experiencing changes in the major environmental parameters, and many studies have addressed the observed changes, climate feedbacks and their consequences (e.g. [18,22,23]). However, there remains an incomplete understanding of the mechanisms responsible for the observed changes on decadal time scales. Figure 1 highlights significant variability of all parameters on a range of time scales shorter than 30 years, occurring in the Arctic since 1948. Trends and variability in most parameters appear to have accelerated relatively recently (since

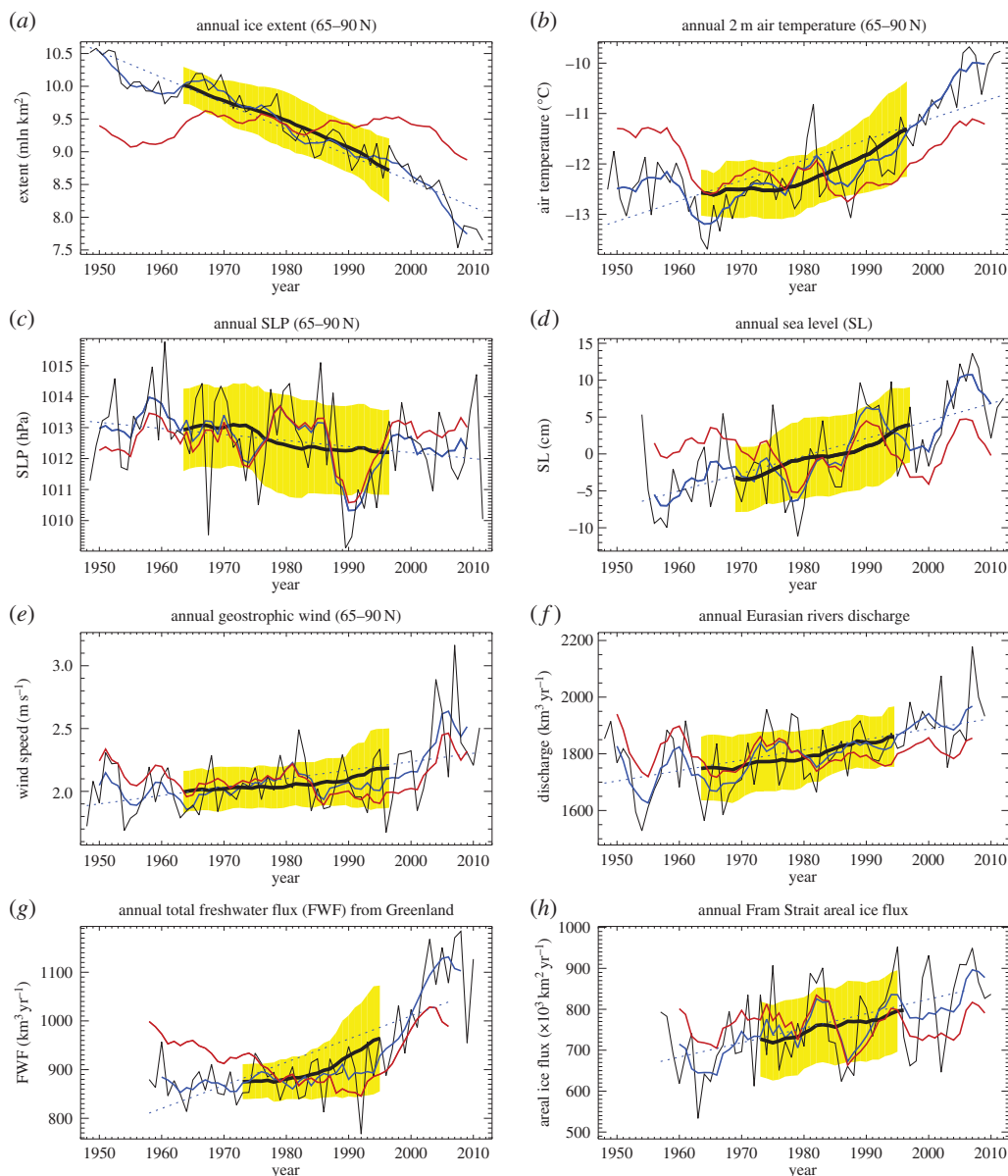


Figure 1. (a) Timeseries of annual ice extent, (b) 2 m air temperature, (c) SLP, (d) sea level, (e) geostrophic wind, (f) Eurasian river runoff, (g) freshwater flux from Greenland and (h) Fram Strait areal ice flux (sources of data are shown in table 1). Thin black line depicts annual, blue line shows 5-year running mean, red line is the detrended 5-year running mean, thick black line shows climatic changes as the 30-year running mean, dotted line is the 1948–2013 linear trend, and the yellow shaded area shows ± 1 s.d. of the 30-year mean series.

approx. 1980s–1990s; figure 1), and already influence climate curves shown in figure 1 (compare observed and detrended timeseries). However, we cannot estimate the present (2014) climate state until we have observations spanning 2000 to 2029. In this sense for example, the fate of Arctic sea ice for at least the next decade remains highly indeterminant despite predictions of a seasonally ice-free Arctic in the near future [24–26]. This problem of uncertainty is well known and is due to short (on climatic time scales) observational records and longer natural variability of the Arctic climate system.

Table 1. Data sources.

data	data description and data source
SLP	four times daily derived from NCAR/NCEP reanalysis product [14]. <i>Reanalysis data provided by the NOAA/OAR/ESRL PSD, Boulder, CO, USA, from their website at http://www.esrl.noaa.gov/psd/</i>
sea ice extent	monthly 25 km grid resolution product from NSIDC derived by Comiso 1999, updated 2012. Bootstrap sea ice concentrations from Nimbus-7 SMMR and DMSP SSM/I-SSMIS (26 October 1978 through 31 December 2011). Boulder, CO, USA: National Snow and Ice Data Center. Digital media
2 m air temperature	monthly values for 1948/01–present. <i>NCEP reanalysis</i> [14]. <i>Derived data provided by the NOAA/OAR/ESRL PSD, Boulder, CO, USA, from their website at http://www.esrl.noaa.gov/psd/</i>
Eurasian rivers discharge	monthly 1935–2011 discharges from Ob', Yenisey, Lena, Severnaya Dvina, Pechora and Kolyma. The data were provided by A. I. Shiklomanov (Water Systems Analysis Group Institute for the Study of Earth, Oceans, and Space, University of New Hampshire, Durham, NH 03824). Data are also available online from the US Geological Survey (http://waterdata.usgs.gov/ak/nwis) and Environment Canada (http://www.wsc.ec.gc.ca/applications/H2O/index-eng.cfm)
sea level	monthly relative sea-level data. Approximately 70 tide-gauge stations in the Barents and Siberian Seas (Kara, Laptev, East Siberian and Chukchi Seas) have recorded sea-level changes from the 1950s through the 2000s. These data are available for model validation at the Permanent Service for Mean Sea-Level archive (http://www.pol.ac.uk/psmsl/pub/nucat.dat) and at the Woods Hole Oceanographic Institution website (http://www.whoi.edu/science/PO/arcticsealevel)
Greenland freshwater flux	monthly freshwater fluxes from Greenland [15]
areal ice flux via Fram Strait	1957–2010 monthly ice area flux from [16] and timeseries received directly from L. H. Smedsrud (2011, personal communication)

(b) Decadal change indices and motivation

Consistent with figure 1, many studies (e.g. [27–30]) indicate Arctic climate variability on a wide range of time scales, with seasonal, decadal and multi-decadal peaks (table 2). This apparent presence of variability at statistically significant frequencies (table 3) suggests potential for predictability of Arctic conditions. There has long been motivation to find basic mechanisms that cause the observed changes in the Arctic atmosphere–ice–ocean system (e.g. [31–38]) at decadal and shorter time scales. Climate indices such as the North Atlantic Oscillation (NAO; e.g. [39,40]), Arctic Ocean Oscillation (AOO) [1–4] and Arctic Oscillation (AO) [41] have been constructed to characterize complex climate processes and explain past variability. However, attempts to predict the NAO and AO cycles were not successful (e.g. [42–44]). This study investigates causes and mechanisms regulating the intensity and duration of Arctic circulation regimes and puts forward an explanation for the well-pronounced decadal changes in the region during the 1948–1996 period and the apparent cessation of the quasi-decadal cycle of circulation regimes after 1996.

Compared to the NAO and AO, the AOO index, defined on the basis of a wind-driven simulated sea surface height field across the Arctic [1], is a more Arctic-centric index. The index is a measure of the intensity and sense (clockwise/anticyclonic or counterclockwise/cyclonic) of the Arctic Ocean wind-driven upper oceanic circulation. Here, we use this index to identify and explain mechanisms regulating changes in Arctic circulation regimes and environmental

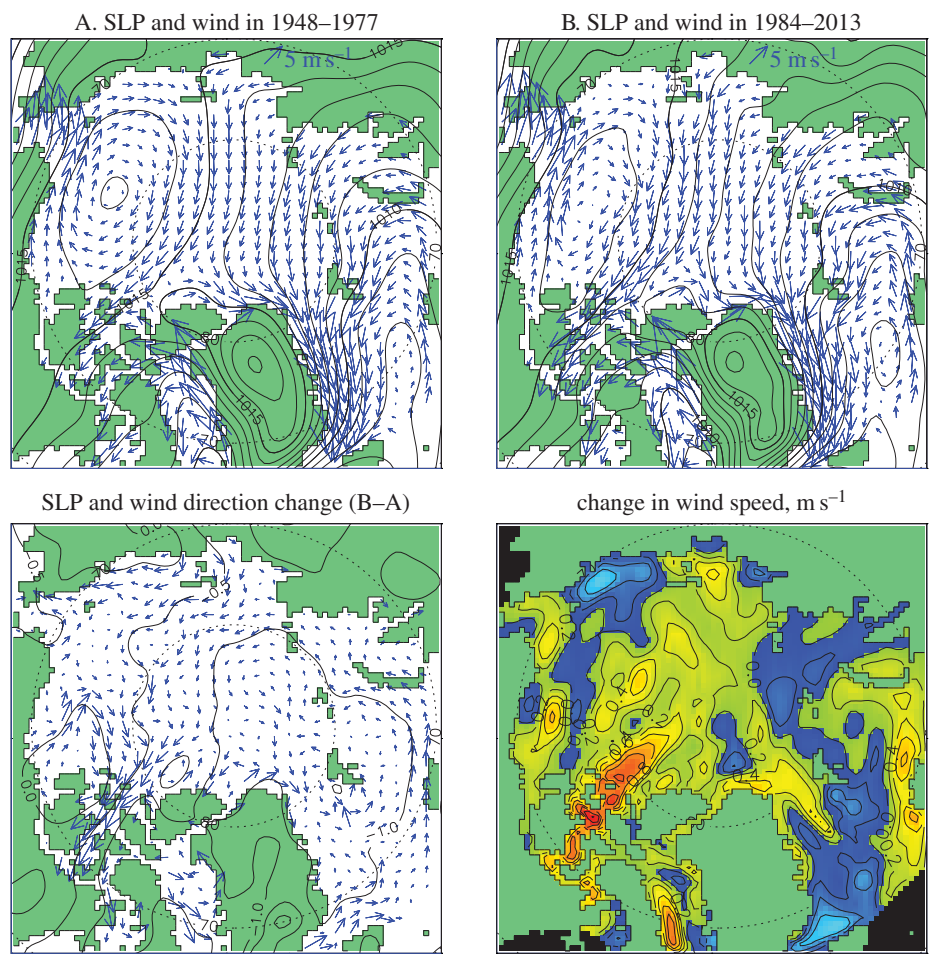


Figure 2. Upper panels show climatic annual SLP (hPa) and geostrophic wind (m s^{-1}) over 1948–1977 (A) and over 1984–2013 (B). Bottom panels show differences in SLP and geostrophic wind direction between B and A climates (left) and change in geostrophic wind speed (right). Reduction of wind speed is depicted in bluish colours.

Table 2. Results of spectral and wavelet analysis, where IE is sea ice extent; T, 2 m air temperature; SLP, sea-level pressure; SL, sea level; R, Eurasian river discharge; F, Fram Strait ice area flux; G, total Greenland freshwater flux; AOO, AOO index. Crosses (×) mean that peak is significant with greater than 50% and less than 90% confidence interval and bold crosses (×) depict peaks with confidence interval greater than 90%.

period (years)	IE	T	SLP	SL	R	F	G	AOO
22–19	×	×	×	×			×	
16–14		×						×
13–11		×			×	×		
10–7	×		×	×		×		×
6–3		×	×	×	×	×		×
2–0.75	×	×			×	×	×	

parameters at decadal time scales. Over 1948–1996, the AOO has been shown to be the most appropriate index to capture variability in key Arctic environmental parameters [2,45,46]. After 1996, many previously established correlations between climate indices and environmental parameters no longer appear to apply. For example, after 1996, despite a low positive or neutral

Table 3. Linear correlation coefficients among 5-year annual running mean climate parameters timeseries for 1960–2009, where IE is sea ice extent; T, 2 m air temperature; SLP, annual mean SLP over the Arctic Ocean; SL, sea level; R, Eurasian river discharge; G, total Greenland freshwater flux; F, Fram Strait ice area flux; AO, Arctic Oscillation index; AOO, AOO index. Bold font depicts correlation coefficients for detrended timeseries shown in figure 1.

	IE	T	SLP	SL	R	G	F	AO	AOO
IE	1.00	−0.94	0.21	−0.82	−0.86	−0.83	−0.81	−0.53	0.42
	1.00	−0.60	−0.21	−0.30	−0.31	−0.40	−0.28	0.07	−0.55
T	−0.94	1.00	−0.05	0.75	0.86	0.92	0.82	0.38	0.53
	−0.60	1.00	0.37	0.09	0.44	0.66	0.36	−0.25	0.61
SLP	−0.82	−0.05	1.00	−0.52	−0.15	0.07	−0.23	−0.88	0.63
	−0.21	0.37	1.00	−0.52	0.15	0.49	−0.02	−0.92	0.76
SL	−0.82	0.75	−0.52	1.00	0.60	0.75	0.67	0.67	0.20
	−0.30	0.09	−0.52	1.00	−0.23	0.33	0.13	0.42	−0.07
R	−0.86	0.86	−0.15	0.60	1.00	0.76	0.66	0.51	0.42
	−0.31	0.44	0.15	−0.23	1.00	0.26	0.00	0.10	0.34
G	−0.83	0.92	0.07	0.75	0.76	1.00	0.63	0.23	0.69
	−0.40	0.66	0.49	0.33	0.26	1.00	0.03	−0.44	0.76
F	−0.81	0.82	−0.23	0.67	0.67	0.63	1.00	0.47	0.21
	−0.28	0.35	−0.02	0.13	0.00	0.03	1.00	0.05	−0.05
AO	−0.53	0.38	−0.88	0.67	0.51	0.23	0.47	1.00	−0.38
	0.07	−0.25	−0.92	0.42	0.10	−0.44	0.05	1.00	−0.67
AOO	0.42	0.53	0.63	0.20	0.42	0.69	0.21	−0.38	1.00
	−0.55	0.61	0.76	−0.07	0.34	0.76	−0.05	−0.67	1.00

AO index, a strong anticyclonic circulation regime (ACCR) has persisted over the Arctic Ocean instead of the more cyclonic circulation as would be expected from previous correlations between the AO index and sea-ice drift direction [47,48]. At the same time, the Arctic experienced continued warming and loss of sea ice extent and volume [49].

In §2, we review the AOO index, updating definitions and results of Proshutinsky *et al.* [1–3] and Dukhovskoy *et al.* [4], explain its physical meaning, and demonstrate that this index accurately reflects changes in the Arctic’s wind-driven circulation and well reproduces existing observational data. In §3, the major anomalies of ACCR and cyclonic circulation regime (CCR) are analysed and decadal variability of the circulation regimes is investigated. We discuss the unusually long ACCR dominating Arctic conditions since 1997 until present in §4 and employ box-model experiments in §5 to speculate on causes of this deviation from expected behaviour. In §6, we summarize and conclude.

2. Arctic Ocean Oscillation index

The AOO was first defined in [1] based on analysis of annual sea surface height fields simulated employing a regional Arctic shallow-water barotropic-coupled ice–ocean model [1,50]. The model was forced by 6-hourly wind stresses derived from NCAR/NCEP atmospheric reanalysis data (table 2) for 1948–1993 and was developed initially to investigate and predict storm surges and tides in the Arctic Ocean [50]. It has been employed since 1992 by the Arctic and Antarctic

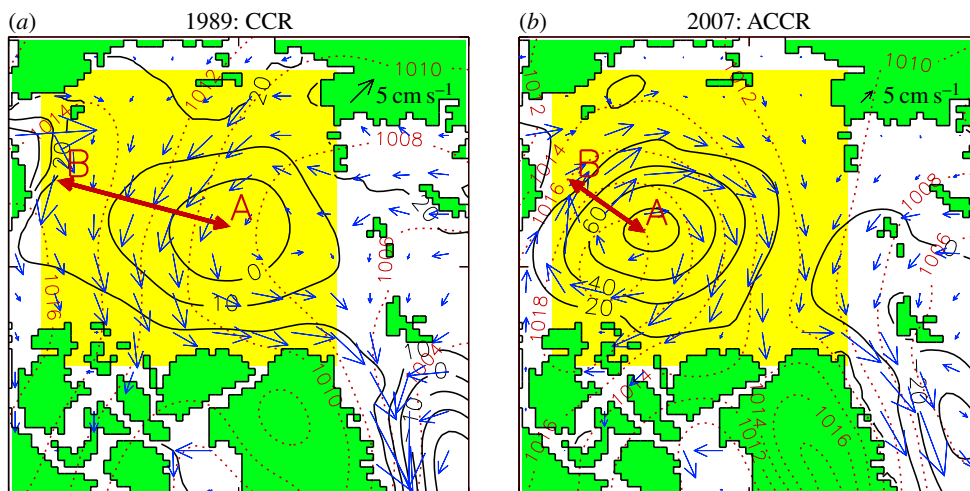


Figure 3. Illustration of the AOO index definition and physical meaning. Annual simulated sea surface heights (cm, contour interval is 10 cm) and ice drift (blue arrows) under CCR (a) and ACCR (b) wind forcing. Red dotted isolines depict SLP (hPa, contour interval is 2 hPa). Yellow box depicts region occupied by Arctic High—major atmospheric circulation feature over the Arctic Ocean. The red arrows in the yellow box show how gradients of sea surface heights (AOO indices) are calculated: differences between sea surface heights in the centre (A) and periphery of the closed circulation (B) are divided by the distance between chosen sea surface height isolines. Note that AOO index is non-dimensional. Anticyclonic circulation has positive sea surface height gradients (AOO indices), while cyclonic circulation has negative gradients (AOO indices).

Research Institute (St Petersburg, Russia) as an operational model to predict sea ice conditions and sea-level variability in the Arctic Ocean. Although only a shallow-water model, full testing and calibration against observed sea-level timeseries along the Siberian coastline and sea-ice drift data from the International Arctic Buoy programme prove that the model well represents the principal patterns of sea ice dynamics, sea-level variability at coastal tide gauges and ocean circulation [51–53].

To calculate the AOO index [1], the location of the extremum in the annual simulated sea level in the Arctic Ocean region is first located and then the last closed sea surface height isoline around that extremum is found to determine the sea surface height horizontal gradient (this number ($\times 10^{-6}$) is non-dimensional and its calculation procedure is illustrated in figure 3). While the magnitude of the gradient is somewhat sensitive to the shape of the last closed isoline, the sign of the gradient is robust. Positive values of these gradients correspond to a raised central sea surface and anticyclonic (clockwise) water circulation (figure 3b). Negative values correspond to lowered central sea surface and cyclonic (counterclockwise) water circulation (figure 3a). Fig. 8 in [1] illustrates the mean annual simulated sea surface heights in the Arctic Ocean for 1948–1993. Figure 4 in this paper shows annual circulation patterns and AOO indices for 1994–2013.

An objective measure of the AOO index timeseries employs empirical orthogonal function analysis (EOF) to derive the major modes of variability for annual sea surface height values for 1948–2013. The first EOF mode of the sea surface height pattern (figure 5e,f) describes 48% of the annual sea surface height variability for this period. Differences in amplitude between the EOF-based AOO and that derived following the manual approach [1,2] are to be expected, because the EOF-based index describes only 48% of variability, while the manually derived AOO measures the full strength of the cyclonic/anticyclonic circulation. Note that these differences change none of the general results here, and the two methods yield indices that are highly correlated ($r = 0.89$) and demonstrate the dominant decadal signal of change in the Arctic Ocean wind-driven circulation regime. The manually derived AOO is analysed here.

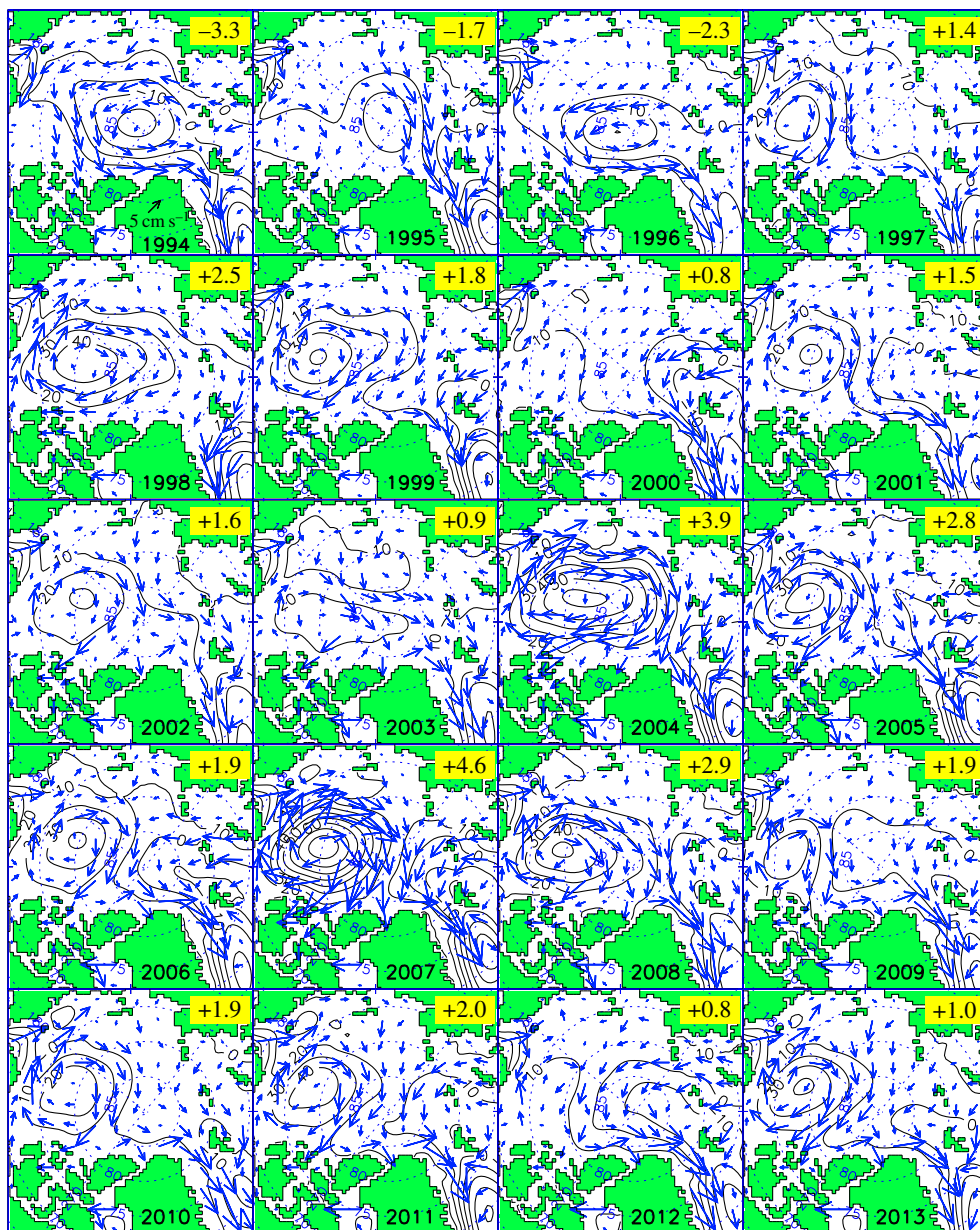


Figure 4. Simulated annual sea surface heights (black contour lines, cm) and wind-forced ice drift component (blue arrows) after 1993. The sea surface height contour interval is 5 cm. The value of the A00 index is shown in the top right corner of all panels. Annual sea surface heights and wind-forced ice drift patterns for 1946–1993 are shown in [1].

3. Arctic Ocean Oscillation variability, circulation regimes and their characteristics

The AOO strongly correlates with many environmental parameters [1,2], characterizing interannual change in the Canada Basin [3], Beaufort Gyre freshwater content [55,56] and variability of freshwater content across the entire Arctic Ocean [57]. During negative AOO years (red bars in figure 5f), the Arctic atmosphere is relatively warm and humid due to frequent penetration of cyclones from the North Atlantic (figure 5a,c) into the central Arctic [1,2]. This

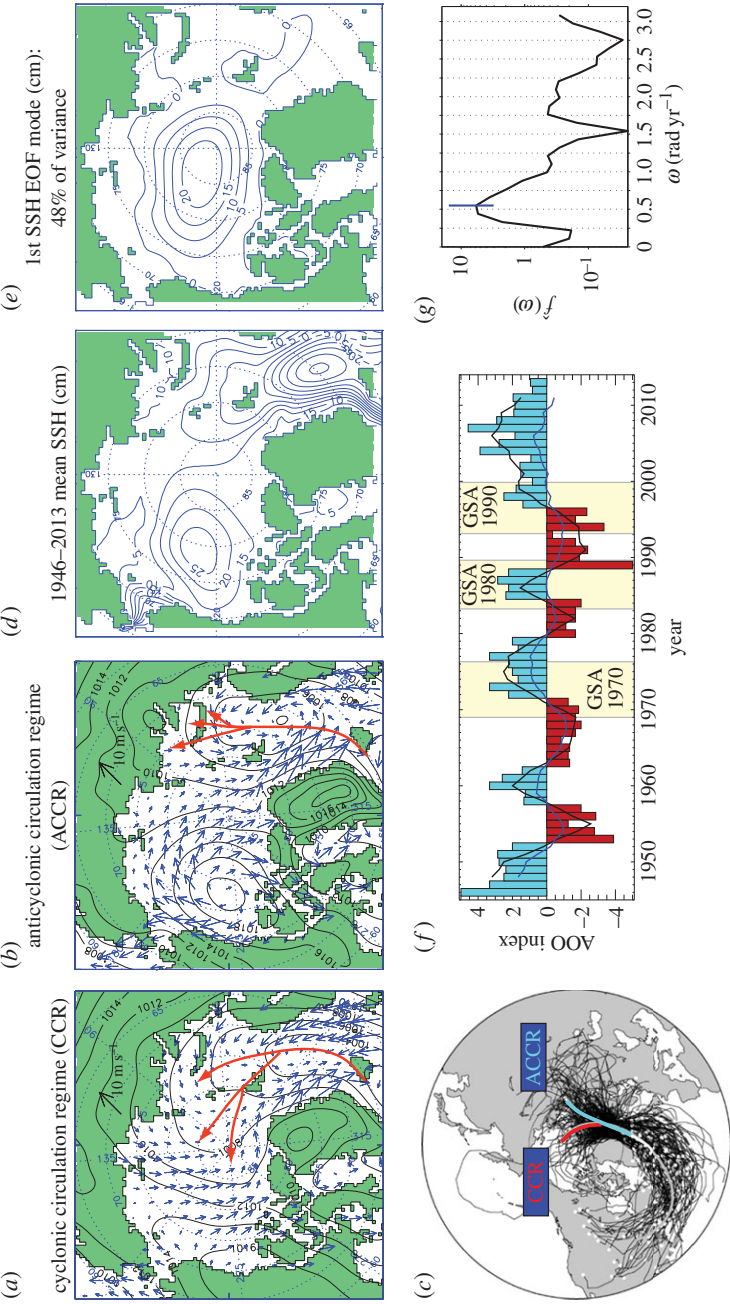


Figure 5. AAO index and circulation regimes. Panels (a) and (b) show typical annual distributions of SLP and surface winds for cyclonic and anticyclonic atmospheric circulation regimes, respectively; red arrows show prevailing cyclone tracks, summarized from panel (c). (c) Greenland Sea cyclone trajectories for 1949–2002 (fig. 2 from [54]). White dots indicate the start of each cyclone trajectory and the white and grey lines show two different paths identified by cluster analysis and well supporting SLP distributions corresponding to circulation regimes shown in (a) and (b), respectively. This means that annual SLP distribution patterns express statistics of cyclone counts—prevailing cyclone trajectories. (d) 1946–2013 mean wind-forced simulated sea surface height (SSH) distribution. (e) First EOF mode (48% of the variance) of the annual sea surface height pattern. (f) AAO indices. The thick black line depicts the 5-year-running-mean timeseries of the AAO index derived manually following the approach of Proshutinsky & Johnson [7]. Positive (blue bars) AAO indices correspond to years with ACCRs and negative (red) bars show cyclonic regimes. Coefficients of the annual first EOF mode of the simulated wind-driven sea surface heights as the 5-year running mean are shown as the blue solid line. GSA years are shown as shaded boxes. (g) Spectra of the AAO calculated from the timeseries of the detrended annual AAO index from 1946 to 2013 using the Tukey window with band width $0.279 \text{ rad yr}^{-1}$. The vertical blue bar denotes the 95% confidence interval of the peak with corresponding frequency $\omega = 0.551 \text{ rad yr}^{-1}$ (11.4 years).

is consistent with results of [54] (figure 5c), where it was shown that cyclones that originate in the North Atlantic are a major source of moisture and heat to the Arctic. These cyclones have two distinct trajectories bringing moisture into the central Arctic or Siberia, depending on the atmospheric circulation regime. In the regime discussed above (hereinafter, CCR), low SLP dominates over the central Arctic Ocean (figure 5a), with winds forcing the sea ice to drift cyclonically (counterclockwise). During a CCR, both exports of sea ice (the solid component of the freshwater flux) and low salinity polar water (liquid freshwater flux) to the Greenland–Iceland–Norwegian (GIN) Sea increase [1,2]. Anomalously high freshwater flux results in freshening of the upper GIN Sea. The Great Salinity Anomalies (GSAs) observed in the North Atlantic Subpolar Region (NASR) in the 1970s, 1980s and 1990s (figure 5f; e.g. [58–60]) correlate well with the history of the advent of CCRs starting several years after the initiation of CCRs over the Arctic Ocean.

By contrast, during positive AOO years (figure 5f, blue bars), the trajectories of North Atlantic cyclones are shifted eastward resulting in fewer cyclones reaching the central Arctic (figure 5b,c). In this climate regime (hereinafter, ACCR), high SLP dominates over the Arctic with anticyclonic winds forcing sea ice to drift clockwise (figure 5b). During a typical ACCR, the Arctic atmosphere is relatively cool and dry [1,2]. Lower-than-normal air temperatures and anticyclonic winds lead to thicker ice and increased ice extent in an ACCR compared with a CCR. During an ACCR, freshwater is accumulated in the Beaufort Gyre of the Canada Basin [3,55] due to Ekman transport convergence, reducing freshwater transport towards the NASR (which includes the Labrador Sea and Nordic Seas). This could lead to reduced stratification of the upper ocean there, promoting deep convection and atmospheric warming [3,4].

Results of Proshutinsky *et al.* [2] showed that the two-climate regime theory characterized by AOO variability explains much of the observed decadal variability of the Arctic Ocean and helps to reconcile the different conclusions between analyses of environment data obtained during different climate regimes. Recently, the AOO index was used in several studies to explain decadal-scale regime shifts in northwest Atlantic shelf ecosystems [61] and understand climatic regulations of growth rates of bivalve shells around Svalbard and in the northwestern Barents Sea (e.g. [62]). Our studies also show that there are significant differences in seasonal variability between years with CCRs and ACCRs [45].

Timeseries of the AOO index from 1948 to 1996 indicate that ACCRs and CCRs alternate at approximately 5- to 7-year intervals with a period of quasi-oscillation of about 10–15 years. In a stark deviation from this pattern, the ACCR started in 1997 has dominated the Arctic over the last 17 years (figure 5f). Of essential importance to our understanding of the evolving Arctic system, and its impact on the global environment, is to discern the causes and consequences of the apparent breakdown in the decadal variability of circulation regimes. Why has the well-pronounced decadal variability observed in 1948–1996 been replaced by relatively weak interannual changes under ACCR conditions after 1996?

4. Conceptual models of Arctic climate variability

Several conceptual models of Arctic climate decadal variability have been introduced since the 1990s [63–66] and recently in [3–6]. Mechanisms for multi-decadal Arctic change have also been studied [29,67], but here we focus in particular on decadal variability. One key element of decadal-change conceptual models is the freshwater flux from the Arctic Ocean to the North Atlantic, and another is the atmospheric heat flux from the North Atlantic to the Arctic. Compelling manifestations of the link between the Arctic and the North Atlantic are: (i) salinity anomalies that originated in the Arctic [58,59,68] and propagated in the subpolar gyre in the 1970s, 1980s and 1990s as GSAs [58,59] and (ii) atmospheric warming and cooling events that are related to the intensity of cyclone activity in the central and eastern Arctic [69,70] and coupled to the Arctic atmospheric circulation [1,41].

To explain the observed decadal changes in Arctic climate, a conceptual hypothesis [3] was formulated. In this hypothesis, the freshwater and heat exchanges between the Arctic Ocean and North Atlantic are self-regulated and their interactions are realized via decadal auto-oscillations.

Based on this work, an idealized multi-box model of the Arctic Ocean and NASR ocean–ice–atmosphere system was developed and employed [4–6] to demonstrate how the system oscillates between an ACCR and a CCR.

A detailed description of the multi-box model is provided in [5], while only a brief outline follows here. The prognostic model consists of two coupled modules: an Arctic Ocean and a NASR module, where each includes a coupled sea ice–ocean box model and an atmospheric model. The sea ice–ocean model consists of a thermodynamic ice model and a mixed layer–pycnocline model, with prognostic variables: water temperature and salinity, mixed layer depth and sea-ice thickness. Ocean heat and salt fluxes are specified at the open boundaries of the two modules (Bering Strait, rivers and North Atlantic), and fluxes between the modules are calculated from simulated temperature, salinity and the density gradient between the Arctic and NASR. The atmosphere in the Arctic model is treated via an energy balance, and surface air temperature in the NASR module is calculated as modelled temperature anomaly superimposed on a climatological daily value. The atmospheric heat flux between NASR and Arctic modules is taken to be proportional to the temperature difference between the two atmospheric modules. The full model is forced by daily solar radiation, wind stress, and river runoff. For validation purposes, results from model experiments reproducing seasonal and decadal variability of the major system parameters were analysed and compared with observations and other models [5].

It is obvious that such a multi-box model has several limitations: it does not include dynamics of the atmosphere, ice and ocean; the modelled system is closed, and there are no influences from the global ocean and atmosphere (this is addressed below in §5); and the model cannot show pathways of freshwater from the Arctic Ocean to the subpolar convective gyres. Nevertheless, the advantage of the box model is its simple formulation that allows one to investigate the basic relationships between components of the studied climate system. This task is not straightforward with coupled atmosphere–ice–ocean models, where extricating the role of internal mechanisms in climate shifts is complicated by too many other factors.

It is important to note that in this multi-box model, regime shifts are controlled by atmospheric heat fluxes from the NASR and freshwater fluxes from the Arctic Ocean. The idealized Arctic–NASR climate system (figure 6) correctly shows that the reduced atmospheric heat advection to the Arctic results in lower-than-normal Arctic atmospheric temperature, higher-than-normal SLP and negative atmospheric vorticity (ACCR). ACCR wind forcing leads to freshwater accumulation in the Arctic Ocean via processes of Ekman convergence and reduction of freshwater flux to the NASR, which increases sea surface salinity in the NASR and promotes intensification of deep convection preconditioned by weaker water column stability [3,71]. During this regime, the NASR releases heat accumulated in the deep ocean layers to the atmosphere leading to a warmer-than-normal atmosphere and to the intensification of cyclogenesis. This atmospheric heat is transported to the Arctic with cyclones and contributes to a positive atmospheric vorticity (CCR) over the Arctic Ocean due to the substantial reduction in the Arctic SLP. This state lasts approximately 5–7 years (half period of oscillation). During the second half of this oscillation (5–7 years), positive atmospheric vorticity (cyclonic winds) forces increased freshwater flux from the Arctic Ocean to the NASR and reduces NASR surface salinity. Consequently, strengthened upper ocean stratification leads to reduced NASR convection, reduced heat flux from the ocean to the atmosphere and from the NASR towards the Arctic Ocean, and the Arctic circulation regime shifts to an ACCR. The ACCR is characterized by predominantly negative vorticity and colder than normal Arctic and NASR atmospheres.

Solutions obtained in the Arctic Ocean and NASR box model [4–6] well reproduced the observed evolution of the major anomalies in the ocean temperature and salinity structure, sea ice volume and freshwater fluxes during ACCRs and CCRs prior to approximately 1997. However, after 1997, the real climate system (figures 1 and 5f) does not behave according to the predictions of the conceptual models described above [3–6]. After 1997, the AOO index is positive (figures 4 and 5f) and exhibits only weak interannual variability associated with the strength of the ACCR with no shift in annual mean conditions from an ACCR to CCR. Similarly,

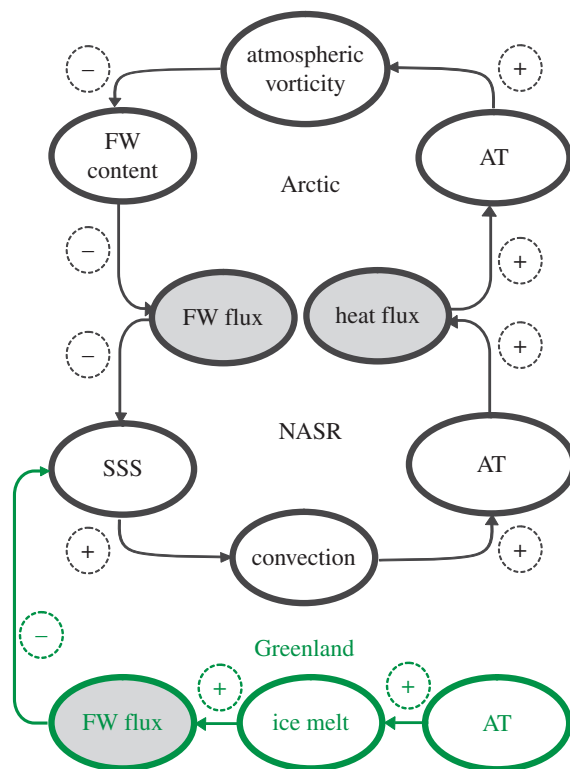


Figure 6. Arctic–NASR idealized atmosphere–ice–ocean climate system with feedbacks regulating decadal variability. A plus sign denotes mechanisms with positive feedback between two cells (i.e. an increase/decrease in one cell causes an increase/decrease in the second), a minus sign denotes mechanisms with negative feedback (i.e. an increase/decrease in one cell causes a decrease/increase in the second). Portions in black represent the idealized behaviour of a closed system under stable climate conditions. Green portions of the schematic represent the influence of FW fluxes from Greenland on parameters of the Arctic–NASR system. For this case, the new system is not closed. AT, SSS and FW stand for air temperature, sea surface salinity and freshwater, respectively.

other conceptual models cited above do not have a mechanism to explain and predict the recent large-scale anticyclonic circulation observed in the Arctic.

5. Increased Greenland freshwater flux as a mechanism for regime shift cessation

Following our conceptual model, the prevailing ACCR should have switched to a CCR in the early 2000s. Instead, the Arctic has been characterized by an ACCR since 1997 at least until 2013 (figure 5f). Results of the Beaufort Gyre Observing System [55,56,72] indicate anomalously high freshwater accumulation (5000 km³ relative to reference salinity of 34.8) in the Beaufort Gyre since the start of the present ACCR in 1997. At the same time, the GIN Sea (as part of the NASR) shows a warming of the deep layers in the 2000s that has been attributed to a cessation of deep convection in the region [73]—conditions that typically develop during a CCR when freshwater is released from the Arctic Ocean. Following our conceptual model, one would expect enhanced deep convection and cyclone formation over the NASR (enabled by reduced freshwater fluxes from the Arctic) would have resulted in a regime shift from ACCR to CCR nearly a decade ago due to freshwater release from the Arctic Ocean. However, analyses of freshwater fluxes through Fram Strait [74,75] purport that the Fram Strait annual mean freshwater flux does not show any large variations since 2002.

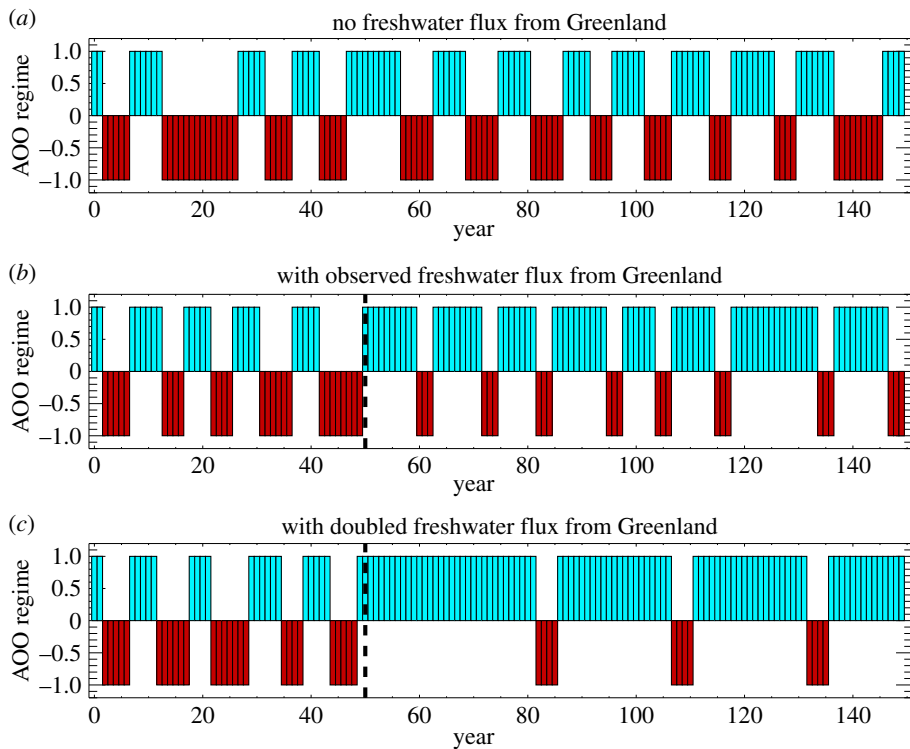


Figure 7. Duration of circulation regimes as simulated by the box model. Blue and red bars indicate ACCR and CCR years, respectively. Panels show model runs without (*a*), with observed (*b*) and doubled (*c*) freshwater flux from Greenland. The black vertical dashed line marks the start of the Greenland freshwater flux addition.

We hypothesize that there is at least one additional factor, neglected in the multi-box model formalism, which has been influencing NASR conditions and has disrupted the auto-oscillatory decadal AAO index variability. The discussed Arctic Ocean–NASR system was viewed as a closed system (the black portions in figure 6) but in recent decades, anomalously warm atmospheric temperatures have led to increased Greenland melt, driving an important external forcing to the Arctic Ocean–NASR system (figure 1g and green portions in figure 6). Recent assessments of freshwater flux from Greenland show that during 1992–2010, this flux to the Arctic Ocean and NASR increased by 36% [15]. We speculate that the excess freshwater advected into the NASR may have significant impact to deep convection (with subsequent atmospheric cooling and reduction of cyclonic activity). The effect would be to impede the decadal oscillations that were a feature of the observations and well represented by our idealized multi-box model (without a contribution to freshwater from Greenland) prior to the 2000s. Figure 6 (green portions) represents the influence of freshwater fluxes from Greenland on other environmental parameters in the Arctic–NASR system via reduction of NASR surface salinity, suppressing deep convection and maintaining a negative atmospheric vorticity (ACCR) over the Arctic.

To examine this central idea, we ran a multi-box model simulation incorporating the additional Greenland freshwater source. After 50 years, an additional freshwater flux is superimposed (following an appropriate seasonal cycle) on the total NASR freshwater input; the annual average of the additional freshwater flux is comparable to the estimated annual inflow from Greenland melt (approx. $46 \text{ km}^3 \text{ yr}^{-1}$ [15]). We then ran a second simulation doubling the freshwater flux anomaly to the NASR box.

The control run with no freshwater flux from Greenland produces quasi-decadal oscillations between ACCRs and CCRs (figure 7a) that are similar to AAO variability (figure 5f) until 2002, while the additional freshwater flux resulted in significantly longer ACCRs (figure 7b),

consistent with observations after 1996. A doubling of the volume flux of freshwater (figure 7c) from Greenland to the NASR (not inconceivable in the future given present warming and Greenland melting trends [76,77]) resulted in extreme changes in the system behaviour with ACCRs dominating for more than three decades, separated by CCRs with 3–4 year duration. These idealized modelling results are consistent with our broad hypothesis that additional freshwater fluxes from Greenland melt may be sufficient to suppress previously established Arctic decadal variability.

In the experiments described above, freshening of the upper NASR plays the key role in disrupting the auto-oscillatory behaviour of the Arctic Ocean–NASR system. An additional freshwater flux from Greenland may impact water column stability and suppress convection in the NASR, reducing air–sea heat fluxes and leading to lower than normal atmospheric temperature in the NASR and Arctic Ocean. Freshening of the upper NASR also acts to maintain the small dynamic height difference between the Arctic Ocean and the NASR (and to reduce the potential for freshwater release from the Arctic Ocean to the NASR) and within the generally accepted view of convection processes in the NASR (e.g. [68,78–80]).

There have been several numerical studies, employing models having a range of complexities and resolutions, to investigate the role of freshwater flux increase from Greenland in the intensity of the Atlantic meridional overturning circulation [81–83]. Results show that depending on model resolution, applied forcing and duration of simulation, the influences of Greenland ice melt on the Atlantic meridional overturning circulation and on the Arctic and NASR conditions differ substantially. In general, however, results of these model studies indicate that freshwater release from Greenland weakens or interrupts decadal variability but mechanisms and processes responsible for decadal changes are not discussed.

6. Future scenario for circulation regimes

Based on the analysis presented here, we speculate that:

- ocean–atmosphere heat fluxes in the NASR vary with Arctic circulation regimes and regulate interactions between the Arctic Ocean and NASR. Ocean to atmosphere heat fluxes in the NASR are larger during ACCRs (compared with CCRs) supporting cyclogenesis and ultimately a regime shift to a CCR;
- the duration of ACCRs and CCRs in a changing climate will be different from those in the twentieth century; a new mode of variability in the Arctic may consist of long-duration ACCRs, separated by relatively short-duration CCRs; and
- the major cause of cessation of decadal variability is the monotonically increasing freshwater flux anomaly from Greenland that began in the mid-1990s.

One important uncertainty in our analysis is that available time series observations are not sufficiently long to verify the hypothesized causality of the Arctic climate variability and Greenland melt (which itself varies on multi-decadal time scales [76,77]). Nevertheless, the data together with idealized modelling and physical context provide compelling evidence that input of fresh Greenland melt to the surface high-latitude regions can interrupt decadal variability of the Arctic–NASR system.

While most predictions of future Arctic climate under global warming are done by extending observed trends in environmental variables, the effects of global warming are likely not to be monotonic due to complex feedbacks and relationships between Arctic environmental parameters, as shown here. We speculate that under the new scenario Greenland would continue melting and supporting longer duration ACCRs, which could result in Arctic cooling accompanied by increased ice extent and thickness—similar to conditions observed in the 1970s. Enhanced upper ocean stratification in the NASR could continue to inhibit heat exchange between the ocean and atmosphere and, subsequently, transport of heat by cyclones to the Arctic. This

would have broad-reaching climate implications by limiting the advection of heat from the mid-latitudes and is therefore a negative feedback on Arctic amplification under global warming.

Data accessibility. All datasets are publicly available and data sources are listed in table 1.

Authors' contributions. A.P. and D.D. formulated hypothesis, designed and conducted experiments. M.-L.T., R.K. and J.B. provided and processed observational data and validated model results. All were involved in paper preparation.

Competing interests. We declare we have no competing interests.

Funding. Support was provided by US National Science Foundation PLR 1313614, 1203720, 1107277 and 0856531 to A.P., PLR-0804017 to D.D. and by the HYCOM consortium (no. N00014-09-1-0587) to D.D.

References

1. Proshutinsky A, Johnson M. 1997 Two circulation regimes of the wind-driven Arctic Ocean. *J. Geophys. Res.* **102**, 12 493–12 514. (doi:10.1029/97JC00738)
2. Proshutinsky A, Polyakov I, Johnson M. 1999 Climate states and variability of Arctic ice and water dynamics during 1946–1997. *Pol. Res.* **18**, 135–142. (doi:10.1111/j.1751-8369.1999.tb00285.x)
3. Proshutinsky A, Bourke RH, McLaughlin FA. 2002 The role of the Beaufort Gyre in Arctic climate variability: seasonal to decadal climate scales. *Geophys. Res. Lett.* **29**, 2100. (doi:10.1029/2002GL015847)
4. Dukhovskoy DS, Johnson M, Proshutinsky A. 2004 Arctic decadal variability: an auto-oscillatory system of heat and fresh water exchange. *Geophys. Res. Lett.* **31**, L03302. (doi:10.1029/2003GL019023)
5. Dukhovskoy DS, Johnson M, Proshutinsky A. 2006 Arctic decadal variability from an idealized atmosphere–ice–ocean model: 1. Model description, calibration, and validation. *J. Geophys. Res.* **111**, C06028. (doi:10.1029/2004JC002821)
6. Dukhovskoy DS, Johnson M, Proshutinsky A. 2006 Arctic decadal variability from an idealized atmosphere–ice–ocean model: 2. Simulation of decadal oscillations. *J. Geophys. Res.* **111**, C06029. (doi:10.1029/2004JC002820)
7. Nansen F. 1902 Oceanography of the North Polar Basin. The Norwegian polar expedition 1893–1896. *Sci. Res.* **3**, 427.
8. Zubov NN, Somov MM. 1940 Ice drift in the northern Arctic Basin. *Probl. Arkt.* **2**, 51–68.
9. Treshnikov AF. 1959 Surface water in the Arctic Basin. *Probl. Arkt.* **7**, 5–14.
10. Coachman LK, Barnes CA. 1963 The movement of Atlantic water in the Arctic Ocean. *Arctic* **16**, 9–16. (doi:10.14430/arctic3517)
11. Aagaard K, Carmack EC. 1994 The Arctic Ocean and climate: a perspective. In *The polar oceans and their role in shaping the global environment* (eds OM Johannessen, RD Muench, JE Overland). Geophysical Monograph Series, vol. 85, pp. 5–20. Washington, DC: American Geophysical Union.
12. Research Council of the National Academies. 2012 *Seasonal to decadal predictions of Arctic sea ice: challenges and strategies*. Washington, DC: National Academies Press.
13. Research Council of the National Academies. 2014 *The Arctic in the Anthropocene: emerging research questions*. Washington, DC: National Academies Press.
14. Kalnay E *et al.* 1996 The NCEP/NCAR 40-year reanalysis project. *Bull. Am. Meteorol. Soc.* **77**, 437–471. (doi:10.1175/1520-0477(1996)077<0437:TNYP>2.0.CO;2)
15. Bamber J, van den Broeke M, Ettema J, Lenaerts J, Rignot E. 2012 Recent large increases in freshwater fluxes from Greenland into the North Atlantic. *Geophys. Res. Lett.* **39**, L19501. (doi:10.1029/2012GL052552)
16. Smedsrud LH, Sirevaag A, Kloster K, Sorteberg A, Sandven S. 2011 Recent wind driven high sea ice area export in the Fram Strait contributes to Arctic sea ice decline. *Cryosphere* **5**, 821–829. (doi:10.5194/tc-5-821-2011)
17. IPCC. 2007 Climate change 2007: the physical science basis. In *Contribution of working group I to the fourth assessment report of the Intergovernmental Panel on Climate Change* (eds S Solomon, D Qin, M Manning, Z Chen, M Marquis, KB Averyt, M Tignor, HL Miller). Cambridge, UK: Cambridge University Press.

18. Overland JE. 2011 Potential Arctic change through climate amplification processes. *Oceanography* **24**, 176–185. (doi:10.5670/oceanog.2011.70)
19. Carton JA, Chepurin GA, Reagan J, Häkkinen S. 2011 Interannual to decadal variability of Atlantic water in the Nordic and adjacent seas. *J. Geophys. Res.* **116**, C11035. (doi:10.1029/2011JC007102)
20. Kwok R, Untersteiner N. 2011 The thinning of Arctic sea ice. *Phys. Today* **64**, 36–41. (doi:10.1063/1.3580491)
21. Jahn A *et al.* 2012 Arctic Ocean freshwater: how robust are model simulations? *J. Geophys. Res.* **117**, C00D16. (doi:10.1029/2012JC007907)
22. Overland JE, Wang M. 2010 Large-scale atmospheric circulation changes are associated with the recent loss of Arctic sea ice. *Tellus* **62A**, 1–9. (doi:10.1111/j.1600-0870.2009.00421.x)
23. Francis JA, Vavrus SJ. 2012 Evidence linking Arctic amplification to extreme weather in mid-latitudes. *Geophys. Res. Lett.* **39**, L06801. (doi:10.1029/2012GL051000)
24. Maslowski W, Clement Kinney J, Higgins M, Roberts A. 2012 The future of Arctic sea ice. *Annu. Rev. Earth Planet. Sci.* **40**, 625–654. (doi:10.1146/annurev-earth-042711-105345)
25. Wang M, Overland JE. 2009 A sea ice free summer Arctic within 30 years? *Geophys. Res. Lett.* **36**, L07502. (doi:10.1029/2009gl037820)
26. Wang M, Overland JE. 2012 A sea ice free summer Arctic within 30 years—an update from CMIP5 models. *Geophys. Res. Lett.* **39**, L18501. (doi:10.1029/2012GL052868)
27. Mysak LA, Venegas SA. 1998 Decadal climate oscillations in the Arctic: a new feedback loop for atmospheric–ice–ocean interactions. *Geophys. Res. Lett.* **25**, 3607–3610. (doi:10.1029/98GL02782)
28. Deser C, Walsh JE, Timlin MS. 2000 Arctic sea ice variability in the context of recent atmospheric circulation trends. *J. Clim.* **13**, 617–633. (doi:10.1175/1520-0442(2000)013<0617:ASIVIT>2.0.CO;2)
29. Polyakov IV, Johnson MA. 2000 Arctic decadal and interdecadal variability. *Geophys. Res. Lett.* **27**, 4097–4100. (doi:10.1029/2000GL011909)
30. Venegas SA, Mysak LA. 2000 Is there a dominant timescale of natural climate variability in the Arctic? *J. Clim.* **13**, 3412–3434. (doi:10.1175/1520-0442(2000)013<3412:ITADTO>2.0.CO;2)
31. Bradley RS, Miller GH. 1972 Recent climatic change and increased glacierization in the eastern Canadian Arctic. *Nature* **237**, 385–387. (doi:10.1038/237385a0)
32. Budyko MI. 1977 *Climate changes*. Baltimore, MD: Waverly.
33. van Loon H, Rogers JC. 1978 The seesaw in winter temperatures between Greenland and Northern Europe. Part I: general description. *Mon. Weather Rev.* **106**, 296–310. (doi:10.1175/1520-0493(1978)106<0296:TSIWTB>2.0.CO;2)
34. Jones PD, Wigley TML, Wright PB. 1986 Global temperature variations between 1861 and 1984. *Nature* **322**, 430–434. (doi:10.1038/322430a0)
35. Wadhams P. 1994 Sea ice thickness changes and their relation to climate. In *Polar oceans and their role in shaping the global environment* (eds OM Johannessen, RD Muench, JE Overland). Geophysical Monograph Series, pp. 337–361. Washington, DC: American Geophysical Union.
36. Dickson R. 1999 All change in the Arctic. *Nature* **397**, 389–391. (doi:10.1038/17018)
37. Vinnikov KY, Robock A, Stouffer RJ, Walsh JE, Parkinson CL, Cavalieri DJ, Mitchell JFB, Garrett D, Zakharov VF. 1999 Global warming and Northern Hemisphere sea ice extent. *J. Clim.* **286**, 1934–1937. (doi:10.1126/science.286.5446.1934)
38. Häkkinen S, Geiger CA. 2000 Simulated low-frequency modes of circulation in the Arctic Ocean. *J. Geophys. Res.* **105**, 6549–6564. (doi:10.1029/2000JC900003)
39. Walker GT. 1924 Correlation in seasonal variation of weather, IX. *Mem. Ind. Met. Dept.* **25**, 275–332.
40. Hurrell JW, Kushnir Y, Ottersen G, Visbeck M (eds). 2003 *The North Atlantic oscillation: climatic significance and environmental impact*. Washington, DC: American Geophysical Union.
41. Thompson DWJ, Wallace JM. 1998 The Arctic oscillation signature in the wintertime geopotential height and temperature fields. *Geophys. Res. Lett.* **25**, 1297–1300. (doi:10.1029/98GL00950)
42. Wunsch C. 1999 The interpretation of short climate records, with comments on the North Atlantic and Southern oscillations. *Bull. Am. Meteorol. Soc.* **80**, 245–255. (doi:10.1175/1520-0477(1999)080<0245:TIOSCR>2.0.CO;2)
43. Stephenson DB, Pavan V, Bojariu R. 2000 Is the North Atlantic oscillation a random walk? *Int. J. Climatol.* **20**, 1–18. (doi:10.1002/(SICI)1097-0088(200001)20:1<1::AID-JOC456>3.0.CO;2-P)

44. Dickson RR, Osborn TJ, Hurrell JW, Meincke J, Blindheim J, Adlandsvik B, Vinje T, Alekseev G, Maslowski W. 2000 The Arctic Ocean response to the North Atlantic oscillation. *J. Clim.* **13**, 2671–2696. (doi:10.1175/1520-0442(2000)013<2671:TAORTT>2.0.CO;2)
45. Polyakov I, Proshutinsky A, Johnson M. 1999 The seasonal cycles in two regimes of Arctic climate. *J. Geophys. Res.* **104**, 25 761–25 788. (doi:10.1029/1999JC900208)
46. Overland JE. 2009 Meteorology of the Beaufort Sea. *J. Geophys. Res.* **114**, C00A07. (doi:10.1029/2008JC004861)
47. Overland JE, Wang M. 2005 The Arctic climate paradox: the recent decrease of the Arctic oscillation. *Geophys. Res. Lett.* **32**, L06701. (doi:10.1029/2004GL021752)
48. Makshtas A, Atkinson D, Kulakov M, Shutilin S, Krishfield A, Proshutinsky A. 2007 Atmospheric forcing validation for modeling the central Arctic. *Geophys. Res. Lett.* **34**, L20706. (doi:10.1029/2007GL031378)
49. Kwok R, Rothrock A. 2009 Decline in Arctic sea ice thickness from submarine and ICESat records: 1958–2008. *Geophys. Res. Lett.* **36**, L15501. (doi:10.1029/2009GL039035)
50. Proshutinsky A. 1993 *Arctic Ocean level variability*. St Petersburg, Russia: Gidrometeoizdat.
51. Proshutinsky A, Ashik I, Dvorkin E, Häkkinen S, Krishfield R, Peltier R. 2004 Secular sea level change in the Russian sector of the Arctic Ocean. *J. Geophys. Res.* **109**, C03042. (doi:10.1029/2003JC002007)
52. Joyce T, Proshutinsky A. 2007 Greenland's island rule and the Arctic Ocean circulation. *J. Mar. Res.* **65**, 639–653. (doi:10.1357/002224007783649439)
53. Timmermans M-L, Proshutinsky A, Krishfield R, Perovich D, Richter-Menge J, Stanton T, Toole J. 2011 Surface freshening in the Arctic Ocean's Eurasian Basin: an apparent consequence of recent change in the wind-driven circulation. *J. Geophys. Res.* **116**, C00D03. (doi:10.1029/2011JC006975)
54. Sortenberg A, Walsh JE. 2008 Seasonal cyclone variability at 70°N and its impact on moisture transport into the Arctic. *Tellus A* **60**, 570–586. (doi:10.1111/j.16000870.2008.00314.x)
55. Proshutinsky A *et al.* 2009 The Beaufort gyre fresh water reservoir: state and variability from observations. *J. Geophys. Res.* **114**, C00A10. (doi:10.1029/2008JC005104)
56. Krishfield RA, Proshutinsky A, Tateyama K, Williams WJ, Carmack EC, McLaughlin FA, Timmermans M-L. 2014 Deterioration of perennial sea ice in the Beaufort Gyre from 2003 to 2012 and its impact on the oceanic freshwater cycle. *J. Geophys. Res. Oceans* **119**, 1271–1305. (doi:10.1002/2013JC008999)
57. Rabe B, Karcher M, Kauker F, Schauer U, Toole JM, Krishfield RA, Pisarev S, Kikuchi T, Su J. 2014 Arctic Ocean basin liquid freshwater storage trend 1992–2012. *Geophys. Res. Lett.* **41**, 961–968. (doi:10.1002/2013GL058121)
58. Dickson RR, Meincke J, Malmberg SA, Lee LJ. 1988 The great salinity anomaly in the northern North Atlantic 1968–1982. *Prog. Oceanogr.* **20**, 103–151. (doi:10.1016/0079-6611(88)90049-3)
59. Belkin IM, Levitus S, Antonov J, Malmberg S-A. 1998 'Great salinity anomalies' in the North Atlantic. *Prog. Oceanogr.* **41**, 1–68. (doi:10.1016/S0079-6611(98)00015-9)
60. Hakkinen S. 1993 Arctic source for great salinity anomaly: a simulation of the Arctic ice ocean system for 1955–1975. *J. Geophys. Res.* **98**, 16 397–16 410. (doi:10.1029/93JC01504)
61. Greene CH *et al.* 2013 Remote climate forcing of decadal-scale regime shifts in Northwest Atlantic shelf ecosystems. *Limnol. Oceanogr.* **58**, 803–816. (doi:10.4319/lo.2013.58.3.0803)
62. Carroll ML, Ambrose WG, Levin BS, Locke WL, Henkes GA, Hop H, Renaud PE. 2011 Pan-Svalbard growth rate variability and environmental regulation in the Arctic bivalve *Serripes groenlandicus*. *J. Mar. Syst.* **88**, 239–251. (doi:10.1016/j.jmarsys.2011.04.010)
63. Ikeda M. 1990 Decadal oscillations of the air–ice–ocean system in the northern hemisphere. *Atmos. Ocean* **28**, 106–139. (doi:10.1080/07055900.1990.9649369)
64. Mysak LA, Power SB. 1992 Sea-ice anomalies in the western Arctic and Greenland-Icelandic Sea and their relation to an interdecadal climate cycle. *Clim. Bull.* **26**, 147–176.
65. Ikeda M, Wang J, Zhao JP. 2001 Hypersensitive decadal oscillations in the Arctic/subarctic climate. *Geophys. Res. Lett.* **28**, 1275–1278. (doi:10.1029/2000GL011773)
66. Goosse H, Selten F, Haarsma R, Opsteegh J. 2002 A mechanism of decadal variability of the sea–ice volume in the Northern Hemisphere. *Clim. Dyn.* **19**, 61–83. (doi:10.1007/s00382-001-0209-5)
67. Wyatt MG, Curry JA. 2013 Role for Eurasian Arctic shelf sea ice in a secularly varying hemispheric climate signal during the 20th century. *Clim. Dyn.* **42**, 2763–2782. (doi:10.1007/s00382-013-1950-2)

68. Aagaard K, Carmack EC. 1989 The role of sea ice and freshwater in the Arctic circulation. *J. Geophys. Res.* **94**, 14 485–14 498. (doi:10.1029/JC094iC10p14485)
69. Serreze MC, Carse F, Barry RG. 1997 Icelandic low cyclone activity: climatological features, linkages with the NAO, and relationships with recent changes in the Northern Hemisphere circulation. *J. Clim.* **10**, 453–464. (doi:10.1175/1520-0442(1997)010<0453:ILCACF>2.0.CO;2)
70. Rigor IG, Colony R, Martin S. 2000 Variations in surface air temperature observations in the Arctic, 1979–97. *J. Clim.* **13**, 896–914. (doi:10.1175/1520-0442(2000)013<0896:VISATO>2.0.CO;2)
71. Malmberg S-A, Jonsson S. 1997 Timing of deep convection in the Greenland and Iceland Seas. *ICES J. Mar. Sci.* **54**, 300–309. (doi:10.1006/jmsc.1997.0221)
72. Proshutinsky A *et al.* 2012 *Ocean. Bull. Am. Meteorol. Soc.* **93**, S142–S147.
73. Somavilla R, Schauer U, Budéus G. 2013 Increasing amount of Arctic Ocean deep waters in the Greenland Sea. *Geophys. Res. Lett.* **40**, 4361–4366. (doi:10.1002/grl.50775)
74. De Steur L, Hansen E, Gerdes R, Karcher M, Fahrbach E, Holfort J. 2009 Freshwater fluxes in the east Greenland current: a decade of observations. *Geophys. Res. Lett.* **36**, L23611. (doi:10.1029/2009GL041278)
75. Mauritzen C *et al.* 2011 Closing the loop—approaches to monitoring the state of the Arctic Mediterranean during the International Polar Year 2007–2008. *Prog. Oceanogr.* **90**, 62–89. (doi:10.1016/j.pocean.2011.02.010)
76. Frauenfeld OW, Knappenberger PC, Michaels PJ. 2011 A reconstruction of annual Greenland ice melt extent, 1785–2009. *J. Geophys. Res.* **116**, D08104. (doi:10.1029/2010JD014918)
77. Kobashi T, Severinghaus JP, Barnola J-M, Kawamura K, Carter T, Nakaegawa T. 2010 Persistent multi-decadal Greenland temperature fluctuation through the last millennium. *Clim. Change* **100**, 733–756. (doi:10.1007/s10584-009-9689-9)
78. Swift JH. 1986 The Arctic waters. In *The Nordic seas* (ed. BG Hurdle), pp. 129–153. New York, NY: Springer. (doi:10.1007/978-1-4615-8035-5_5)
79. Alekseev GV, Ivanov VV, Korablev AA. 1994 Interannual variability of the thermohaline structure in the convective gyre of the Greenland Sea. In *Polar oceans and their role in shaping the global environment* (eds OM Johannessen, RD Muench, JE Overland). Geophysical Monograph Series, pp. 485–496. Washington, DC: American Geophysical Union.
80. Pawlowicz A. 1995 A note on seasonal cycles of temperature and salinity in the upper waters of the Greenland Sea Gyre from historical data. *J. Geophys. Res.* **100**, 4715–4726. (doi:10.1029/94JC02014)
81. Driesschaert E, Fichet T, Goosse H, Huybrechts P, Janssens I, Mouchet A, Munhoven G, Brovkin V, Weber SL. 2007 Modeling the influence of Greenland ice sheet melting on the Atlantic meridional overturning circulation during the next millennia. *Geophys. Res. Lett.* **34**, L10707. (doi:10.1029/2007GL029516)
82. Swingedouw D, Rodehacke CB, Behrens E, Menary M, Olsen SM, Gao Y, Mikolajewicz U, Mignot J, Biastoch A. 2013 Decadal fingerprints of freshwater discharge around Greenland in a multi-model ensemble. *Clim. Dyn.* **41**, 695–720. (doi:10.1007/s00382-012-1479-9)
83. Marsh R, Desbruyères D, Bamber JL, de Cuevas BA, Coward AC, Aksenov Y. 2010 Short-term impacts of enhanced Greenland freshwater fluxes in an eddy-permitting ocean model. *Ocean Sci.* **6**, 749–760. (doi:10.5194/os-6-749-2010)